

INCREASED TEMPERATURE IN URBAN GROUND AS SOURCE OF SUSTAINABLE ENERGY

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ABSTRACT

Densely urbanized areas are characterized by special microclimatic conditions with typically elevated temperatures in comparison with the rural surrounding. This phenomenon is known as the urban heat island (UHI) effect, but not restricted exclusively to the atmosphere. We also find significant warming of the urban subsurface and shallow groundwater bodies. Here, main sources of heat are elevated ground surface temperatures, direct thermal exploitation of aquifers and heat losses from buildings and other infrastructure. By measuring the shallow groundwater temperature in several European cities, we identify that heat sources and associated transport processes interact at multiple spatial and temporal scales. The intensity of a subsurface UHI can reach the values of above 4 K in city centres with hotspots featuring temperatures up to +20°C. In comparison with atmospheric UHIs, subsurface UHIs represent long-term accumulations of heat in a relatively sluggish environment. This potentially impairs urban groundwater quality and permanently influences subsurface ecosystems. From another point of view, however, these thermal anomalies can also be seen as hidden large-scale batteries that constitute a source of shallow geothermal energy. Based on our measurements, data surveys and estimated physical ground properties, it is possible to estimate the theoretical geothermal potential of the urban groundwater bodies beneath the studied cities. For instance, by decreasing the elevated temperature of the shallow aquifer in Cologne, Germany, by only 2 K, the obtained energy could supply the space-heating demand of the entire city for at least 2.5 years. In the city of Karlsruhe, it is estimated that about 30% of annual heating demand could be sustainably supplied by tapping the anthropogenic heat loss in the urban aquifer. These results reveal the attractive potential of heated urban ground as energy reservoir and storage, which is in place at many places worldwide but so far not integrated in any city energy plans.

Keywords: energy waste, shallow geothermal energy, sustainability, urban aquifers.

1 INTRODUCTION

Increased temperatures beneath cities have been reported in several studies in northern countries. Local surveys in North America (e.g., Ferguson and Woodbury [1], Taylor and Stefan [2], Eggleston and McCoy [3]), Europe (e.g., Balke [4], Menberg *et al.* [5], Müller *et al.* [6], García-Gil *et al.* [7], Epting and Huggenberger [8]), and Asia (e.g., Taniguchi *et al.* [9], Yamano [10], Liu *et al.* [11], Zhan *et al.* [12]) have shown regional thermal anomalies of around 2–5 K in comparison to the undisturbed surrounding. In these subsurface urban heat islands (SUHIs), heat from anthropogenic sources propagates into the urban subsurface increasing the temperature and changing the thermal conditions in shallow ground and groundwater (Fig. 1). The regionally increased temperature is also called SUHI intensity.

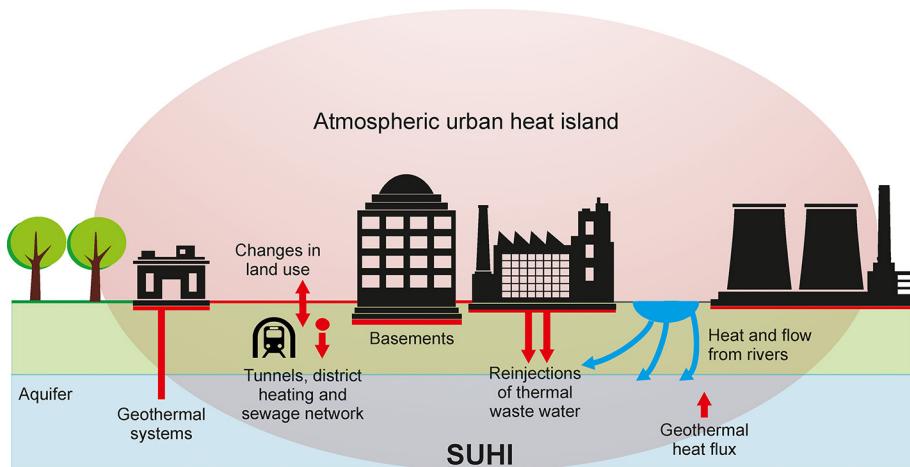


Figure 1: Representation of interacting processes within a subsurface urban heat island (SUHI).

In Europe, one of the first comprehensive studies was presented by Balke [4], who investigated the temperature regime in the shallow aquifer of Cologne, Germany. This was done by temperature measurements in multiple groundwater wells in the city. By constructing contour plots at 20 m depth, a SUHI reaching an intensity of 5 K was resolved beneath the city centre. In similar detail, Hötzl and Makurat [13] investigated the shallow thermal conditions beneath the city of Karlsruhe in 1981. It was shown that the SUHI intensity beneath the city centre reached at least 3 K, and that horizontal groundwater flow promotes the downstream transport of heat.

While more recent studies focused on the environmental impact of SUHIs (e.g Müller *et al.* [6], Bonte *et al.* [14], Briemann *et al.* [15]), others have studied their associated geothermal potential (e.g., Allen *et al.* [16], Arola and Korkka-Niemi [17], Sinnathamby *et al.* [18]). Zhu *et al.* [19], Zhang *et al.* [20] for instance evaluated this potential in several cities including London, Prague, Cologne, Winnipeg and Beijing. The results indicate that by using geothermal systems for decreasing the SUHI intensity, the amount of energy extracted is several times higher than the annual heating demand in these cities. Additionally, the elevated temperatures are crucial for adjusting and enforcing environmental regulations (Hähnlein *et al.* [21], Vienken *et al.* [22]), and they are relevant for technological issues related to conventional shallow geothermal systems (e.g., Banks [23], Rybach and Eugster [24]).

Extracting this shallow geothermal potential requires a comprehensive underground management of the energy content since an unbalanced exploitation scheme progressively degrades the low-enthalpy energy resources (e.g., Rivera *et al.* [25]). Possible ways around are implementation of centralized district heating and/or cooling networks for distributing locally unbalanced energy loads (e.g. Arola and Korkka-Niemi [17], Zhang *et al.* [20]) or direct heat sources. In such “smart energy” management, the urban underground can be conceived as a reservoir that buffers the spatial and seasonal variability of the energy demand.

This work presents a description of the spatial distribution of SUHIs in several German cities including Cologne, Berlin, Munich and Karlsruhe. Finally, the

associated energy content is estimated and compared with case-specific space-heating demands.

2 MATERIALS AND METHODS

2.1 Study cases and available data

The spatial distribution of shallow groundwater temperatures (GWT) was estimated in four German cities. The study cases (Table 1) cover a wide range of population densities with specific patterns in the subsurface thermal alteration. The measured surface air temperatures (SAT) reflect a moderate climate, and the dominant sedimentary deposits are likely sensitive to intensified downward heat fluxes.

The spatial distribution of GWT is obtained from a fairly large number of observation wells that the respective municipalities use for monitoring groundwater quality among other parameters. The wells are not only located within the urban areas but also in the suburban surroundings, which is fundamental in estimating the relative intensity of the subsurface urban heat islands.

In Berlin, these measurements were carried out by the Senate Department for Urban Development and the Environment in 2010 at 0 m above sea level (asl), corresponding to a depth of about 30–60 m below ground level [26]. In Munich, a measurement campaign was conducted in 1983 on a dense network of observation wells along the subway system [27]. In Karlsruhe in turn, the shallow GWT is daily recorded by the Public Works Service through data loggers installed at depths of about 9–12 m. Due to the typical seasonal variability of the temperature data at these depths, the arithmetic annual mean for the year 2012 is used for the present analysis. Finally in Cologne, this study uses the data from Zhu *et al.* [19] that describes the spatial distribution of groundwater temperature in 2009. In all cases, the spatially distributed temperatures are visualized and interpolated via kriging in GIS (ESRI® ArcInfo™ 10.0).

Table1: Characteristics of the studied German cities. The annual mean air temperature values are given for the time period 1961–1990.

Parameter	Berlin	Munich	Cologne	Karlsruhe
Mean altitude (masl) ^(a)	75	518	53	115
Population (pop.) ^(a)	3'421'829	1'407'836	1'036'253	299'103
Pop. density (pop./km ²) ^(a)	3838	4531	2557	1724
Total area (km ²) ^(a)	892	311	405	173
Annual mean SAT (°C) ^(b)	8.9	9.2	10	10.7
Shallow aquifer material	Gravel, sand, silt, clay, till ^(c)	Gravel, sand ^(d)	Gravel, sand ^(e)	Sand, gravel ^(f)
Measurement depth (m.b.g.l) ^(g)	20	2–20	15	9–12
Observation wells ^(g)	123	492	52	82

^(a)Federal Statistical Office (2014); ^(b)German Weather Service (2014); ^(c)Hannappel and Limberg [28]; ^(d)Seiler [29], Zosseder [30]; ^(e)Zhu *et al.* [19]; ^(f)Geyer and Gwinner [31];

^(g)Menberg *et al.* [5].

Table 2: Hydrogeological and thermo-physical parameters of the studied cities.

City	Aquifer thickness [m] min – max (mean)	Porosity [-] ^(e) min – max (mean)	Volumetric heat capacity of solid [MJ m ⁻³ K ⁻¹] min – max (mean)
Berlin	30–50 (40) ^(a)	0.15–0.25 (0.20)	2.0–2.2 (2.1) ^(f)
Munich	5–20 (10) ^(b)	0.20–0.25 (0.23)	2.3–2.5 (2.4) ^(f)
Cologne	10–30 (20) ^(c)	0.15–0.25 (0.20)	2.1–2.2 (2.2) ^(f)
Karlsruhe	10–50 (30) ^(d)	0.15–0.25 (0.20)	2.2–2.6 (2.4) ^(f)

Hydrogeological data and values for porosity and thermal properties taken from: ^(a)Limberg et al. and Thierbach [33]; ^(b)Kerl et al. [34]; ^(c)Zhu et al. [19]; ^(d)Schafer et al. [35]; ^(e)Prinz et al. And Strauss [36]; ^(f)VDI [32].

2.2 Geothermal potential of subsurface urban heat islands

The volumetric heat capacity of the porous medium C_m (kJ m⁻³ K⁻¹) can be used for estimating the energy yield due to a change of temperature ΔT (K) in a given reference volume V (m³). This approach was followed by Zhu *et al.* [19] to quantify the theoretical geothermal potential in urban aquifers with elevated GWT. Provided that the thermo-physical and hydrogeological parameters of the aquifer are known, the heat content Q (kJ) can be calculated by:

$$Q = C_m \Delta T V = [n C_w + (1 - n) C_s] \Delta T V \quad (1)$$

where C_m is obtained assuming a fully saturated medium in which volumetric energy contributions from water C_w and from solids C_s , are weighted according to the medium porosity n . This parameter and the aquifer thickness (required to calculate the volume V) are listed in Table 2 for the studied cities.

The parameters included in eqn. (1) considerably vary in space. In order to cope with this geological heterogeneity, parameter ranges and a respective mean value were defined based on hydrogeological data from literature (Table 2). For the heat capacity of water a fixed value of 4150 kJ m⁻³ K⁻¹ was taken [32].

Following Zhu *et al.* [19], the temperature reduction is set to $\Delta T = 2–6$ K, with a mean value of $\Delta T = 4$ K (oriented at SUHI intensity). The calculated energy content Q is contrasted with the space-heating demand of the individual cities. The latter is estimated by the average living space and the average annual unit heating demand of 50 kWh m⁻² [19]. The average living space areas are taken from Timm [37] for different regions in Germany as follows: Berlin 40 m², Munich 44 m², Cologne 42 m² and Karlsruhe 43 m².

3 RESULTS AND DISCUSSION

3.1 Spatial distribution of groundwater temperatures

The interpolated GWT fields for each city are shown Fig. 2. They exhibit a common pattern in SUHIs with closed isotherms and generally maximum temperatures around the city centres. The observed SUHI intensities decrease towards suburban or rural areas, where anthropogenic affectation is less intense. In these areas, the shallow GWT ranges between 8°C and 11°C in the studied cities mirroring the annual mean SAT indicated in Table 1. The

relative large-scale thermal anomalies observed in the city downtowns are likely associated with long-term and extensive changes in land use. For the largest considered cities, the temperatures vary between 13°C–18°C. In Karlsruhe, however, the highest GWT is observed in an industrial area (western part of the city) where important reinjections of thermal waste water are operated [38].

Another common feature in SUHIs is their spatial heterogeneity which is partially caused by the spatial density and depth of the temperature measurements. This is more evident in Munich where the dense network of observation wells allows to identify hot spots with temperature of up to 20°C nearby building cellars and at the same time, low temperatures close to the background values in green spaces in the inner-city [27]. This heterogeneity is also enhanced by the variety of heat sources within a typical urban area including subway systems, district heating networks, shallow geothermal systems, heat losses from basements in buildings and increased surface and air temperatures [1–2] and [38].

3.2 The estimated theoretical geothermal potential

With the parameters ranges listed in Table 2, the theoretical heat content Q can be estimated for each city. The corresponding minimum, maximum and mean values are shown

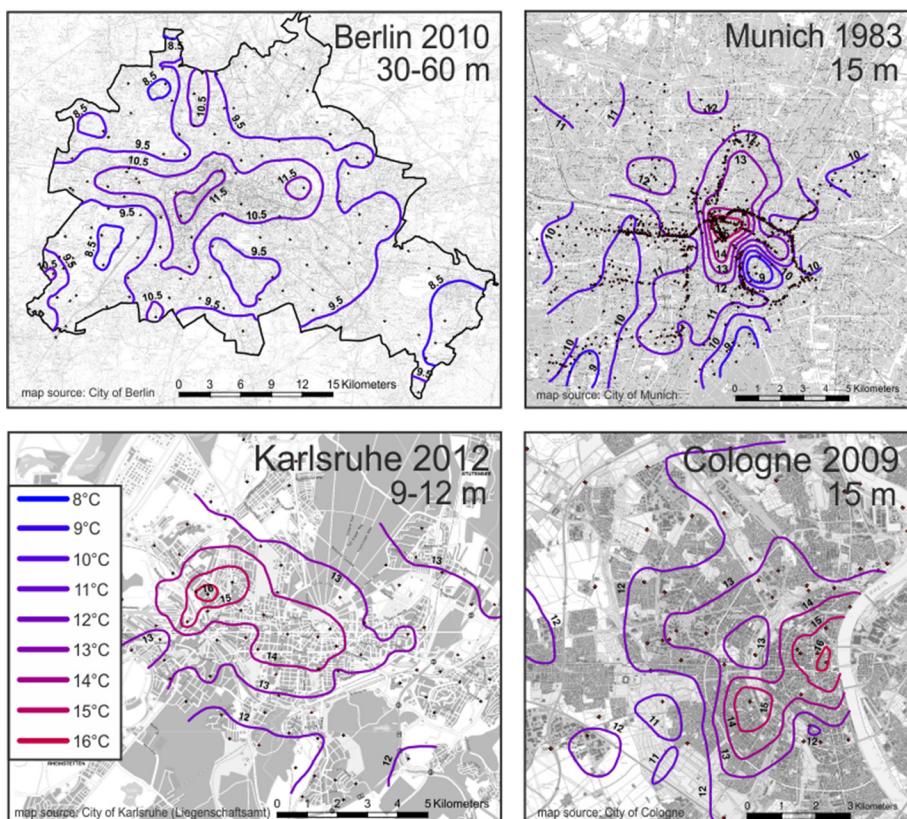


Figure 2: Isotherm maps of groundwater temperature under the studied cities. The dots represent the location of the observation wells [5].

Table 3: Geothermal potential of the studied cities.

City	Potential underground heat content [$\times 10^{12}$ J km $^{-2}$] min – max (mean)	Space-heating demand [$\times 10^{12}$ J km $^{-2}$ year $^{-1}$]	Capacity for space-heating [years] min – max (mean)
Berlin	155–874 (440)	27.8	5.6–31.4 (15.8)
Munich	26.7–350 (112)	20.5	1.3–17.1 (5.5)
Cologne	48.2–484 (204)	19.2	2.5–25.1 (10.6)
Karlsruhe	46.5–806 (301)	36.4	1.3–22.1 (8.3)

in Table 3. Given the similarity in hydro-physical parameters and groundwater warming between the different cities, the estimated potential only varies within one order of magnitude. For comparison, Table 3 also includes the area-averaged space-heating demand. It is shown that the underground stored energy could cover this demand in the individual cities for at least 1.3–5.6 years while more optimistic estimations yield capacities of up to 17.1–31.4 years. These assessments can be judged as conservative since only the actual heat content in the urban aquifer is considered excluding the continuous power input from the associated anthropogenic sources. In fact, provided an appropriate exploitation technology, this continuous replenishment of thermal energy can supply part of the heating demand in a sustainable way. For the city of Karlsruhe for instance, by tapping only the annual anthropogenic heat loss, about 20%–30% of the annual space-heating demand could be covered sustainably [39]. However, these estimations of the geothermal potential are based on a lumped parameterization, which neglects the high variability of heat sources and subsurface hydrogeological conditions. These factors together with technological issues, will ultimately determine the site-specific heat gain from the urban aquifers.

4 CONCLUSIONS

The presented study cases exhibit the increased temperature in the shallow urban subsurface of several German cities. The thermal anomalies have a strong spatial heterogeneity with observed persistent values of up to 16°C. Local temperatures however can reach up to 20°C near dominant heat sources yielding groundwater warming by more than 6 K in comparison to background conditions. Several anthropogenic sources such as heat losses from buildings and other infrastructure, direct geothermal use of aquifers and changes in land cover drive the long-term and multi-scale thermal alteration in the studied urban aquifers.

The accumulated heat content is substantial and urban aquifers can act as attractive and easily accessible reservoirs and also storage for thermal energy and waste heat. The presented estimation of the individual space-heating capacities reveals that the theoretical geothermal potential could fulfil the residential heating demand in all studied cities for at least several years, and assuming more optimistic values even for decades. Considering the permanent anthropogenic heat input from the urban structures into the shallow aquifer, the heating demand of the cities could also be partly satisfied for a longer period. However, to meet the requirements for such a sustainable underground management, a more detailed examination of the transient behaviour of these subsurface urban heat islands with geothermal usage is still needed.

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